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# Understanding coralline algal responses to ocean acidification: Metaanalysis and synthesis

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9 Understanding coralline algal responses to ocean acidification: meta-analysis and synthesis

10 Running title: Coralline algae and ocean acidification

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27

28

29 **Abstract:**

30 Ocean acidification (OA) is a major threat to the persistence of biogenic reefs throughout the  
31 world's oceans. Coralline algae are comprised of high magnesium calcite, and have long been  
32 considered one of the most susceptible taxa to the negative impacts of OA. We summarise  
33 these impacts and explore causes of variability in coralline algal responses using a  
34 review/qualitative assessment of all relevant literature, meta-analysis, quantitative assessment  
35 of critical responses, and a discussion of physiological mechanisms and directions for future  
36 research. We find that most coralline algae experienced reduced abundance, calcification  
37 rates, recruitment rates, and declines in pH within the site of calcification in laboratory  
38 experiments simulating ocean acidification or at naturally elevated CO<sub>2</sub> sites. There were no  
39 other consistent physiological responses of coralline algae to simulated OA (e.g. photo-  
40 physiology, mineralogy and survival). Calcification/growth were the most frequently  
41 measured parameters in coralline algal ocean acidification research, and our meta-analyses  
42 revealed greater declines in seawater pH were associated with significant decreases in  
43 calcification in adults and similar but non-significant trends for juveniles. Adults from the  
44 family Mesophyllumaceae also tended to be more robust to OA, though there was insufficient  
45 data to test similar trends for juveniles. OA was the dominant driver in the majority of  
46 laboratory experiments where other local or global drivers were assessed. The interaction  
47 between OA and any other single driver was often additive, though factors that changed pH at  
48 the surface of coralline algae (light, water motion, epiphytes) acted antagonistically or  
49 synergistically with OA more than any other drivers. With advances in experimental design  
50 and methodological techniques, we now understand that the physiology of coralline algal  
51 calcification largely dictates their responses to OA. However, significant challenges remain,  
52 including improving the geographic and life-history spread of research effort and a need for  
53 holistic assessments of physiology.

54 **Keywords:** Climate change, ocean acidification, coralline algae, CCA. Rhodoliths, maerl,  
55 calcification, meta-analysis

## 56 **Introduction:**

57 Ocean acidification (OA) is the absorption of anthropogenically-derived CO<sub>2</sub> by the surface  
58 seawaters of the world's oceans (Caldeira and Wickett 2003). OA alters seawater chemistry  
59 as follows: hydrogen ions (H<sup>+</sup>) and carbon dioxide (CO<sub>2</sub>) increase, dissolved inorganic  
60 carbon (DIC) and bicarbonate (HCO<sub>3</sub><sup>-</sup>) increase slightly, and carbonate (CO<sub>3</sub><sup>2-</sup>)  
61 concentrations and the calcium carbonate saturation states ( $\Omega$ ) decrease (Feely et al. 2004).  
62 These changes in seawater chemistry have physiological repercussions for many marine  
63 organisms, particularly those that calcify or photosynthesize (Harvey et al. 2013; Kroeker et  
64 al. 2013).

65 Coralline algae are important foundation species in shallow ecosystems from the  
66 tropics to the poles, both forming and cementing reefs, acting as settlement substrates and  
67 nurseries for marine invertebrates, and providing coastal sediments (Nelson 2009; Milazzo et  
68 al. 2014; Cornwall et al. 2021). However, as calcifying species, they are particularly sensitive  
69 to ocean acidification (McCoy and Kamenos 2015; Martin and Hall-Spencer 2017). Earlier  
70 laboratory experiments indicated that the complex changes in seawater carbonate chemistry  
71 occurring under ocean acidification would impact coralline algae through reduced  
72 calcification rates, sizes and/or numbers of recruits, and observations along natural gradients  
73 showed that coralline algae decline in abundance as CO<sub>2</sub> levels in seawater rise (Anthony et  
74 al. 2008; Hall-Spencer et al. 2008; Jokiel et al. 2008; Kuffner et al. 2008; Martin et al. 2008;  
75 Martin and Gattuso 2009). While coralline algae are one of the most sensitive taxa to ocean  
76 acidification (Kroeker et al. 2013), there is considerable variability between different species'  
77 responses (Peña et al. 2021a). This variability could be related to differences in mineralogy,  
78 the ability to regulate favourable carbonate chemistry at the site of calcification, and  
79 morphology (e.g. thallus thickness or geniculate versus non-geniculate) (Nash et al. 2013;  
80 McCoy and Pfister 2014; Cornwall et al. 2017; Barner et al. 2018). Earlier observations were  
81 that ocean acidification could increase mortality through bleaching that could be exacerbated  
82 by warming (Anthony et al. 2008; Martin and Gattuso 2009), as well as change the photo-  
83 physiology (Borowitzka 1981) or mineralogy of these important foundation species (Ries et  
84 al. 2009; Ries 2011).

85 To more accurately predict the future of reefs where coralline algae play essential  
86 ecological roles, the scientific community needs to better understand the drivers of variability  
87 in response to ocean acidification caused by phylogeny, geography and physiology. Earlier

88 reviews on the topic (Hurd et al. 2009; Nelson 2009; Hofmann and Bischof 2014; McCoy and  
89 Kamenos 2015) provided good starting points for the field of research but were too premature  
90 to assess generalisable results properly, the mechanisms responsible, or get accurate estimates  
91 of the magnitude of the effects. Thirteen years from the initial burst of publications and 6  
92 years from the last specialised review, we critically re-evaluate the literature. Here, we aim to  
93 both collate the knowledge of coralline algal responses to ocean acidification and provide  
94 numerical estimates of its effects using meta-analyses and predictive models. The manuscript  
95 is split into 1) qualitative review, 2) meta-analyses, 3) projected declines in cover,  
96 recruitment and calcification, and 4) coralline algal calcification physiology under ocean  
97 acidification and frontiers in future research.

98

## 99 **Methods:**

### 100 *Qualitative analysis:*

101 We searched Web of Science with different combinations of coralline algal terms and “ocean  
102 acidification”. The different coralline algal terms were “coralline algae”, “calcifying algae”  
103 and “CCA”. The list of suitable publications was then cross-checked against the OAICC  
104 database on Pangaea and the list used in Kroeker et al. (2013). This search was completed on  
105 December 1<sup>st</sup>, 2020 and resulted in 298 papers (see Table S1). We then compiled the main  
106 effects of ocean acidification on the coralline algae from each suitable paper, the main effects  
107 of any other assessed driver (e.g. temperature or seasonality), the directionality of the  
108 interaction between OA and the other driver (antagonistic, synergistic, additive, or one or  
109 both and their interaction are unmeasurable Crain et al. 2008), and which driver was the  
110 putative “dominant” driver. We used an approach where four of our author team reviewed the  
111 findings of every study to reduce bias and misinterpretation. In some instances, we also  
112 assessed each study using an  $\alpha = 0.05$ , rather than the reduced alpha used in some  
113 research as corrections for multiple analyses are not optimal. Concerning the main effects of  
114 OA, we classified responses as “net dissolution or complete mortality or lack of  
115 recruitment/cover”, statistically significant negative, statistically significant parabolic  
116 (including both positive and negative parabolic), no measurable effect, statistically significant  
117 positive effect. Hereafter, we refer to these scores as dissolution or complete removal,  
118 negative, parabolic, no measurable effect, and positive. By dominant driver, we qualitatively  
119 refer to the one with the most considerable effect that can assist future decisions regarding

120 experimental design in multi-driver experiments (sensu Boyd et al. 2015; Boyd et al. 2018).  
121 We refer to these (usually) environmental factors as “drivers” because their effects did not  
122 always manifest in a negative direction. We found 392 responses to OA that could be scored  
123 in this manner. Here, we refer to “responses” as all combined treatments in any one  
124 experiment per species.

125

126 *Meta-analysis:*

127 During this process, it became evident that adult calcification/linear extension/growth and  
128 juvenile growth/linear extension were the most commonly recorded parameters. Because of  
129 their importance as overall indicators of “success” under these OA experiments, we chose to  
130 conduct meta-analyses to further explore patterns in these responses. We used the same  
131 papers identified above. To obtain the calcification and growth data, we extracted means of  
132 calcification rates and their associated error whenever they were listed. When data were not  
133 deposited freely online those data were extracted from figures within publications using the  
134 software Datathief (<http://www.datathief.org>).

135 Seawater carbonate chemistry (pH<sub>T</sub>, total alkalinity, dissolved inorganic carbon),  
136 salinities, and temperatures were extracted or calculated along with the calcification data in  
137 instances where they were not given. Seawater carbonate chemistry was recalculated in some  
138 studies where inconsistencies were found between pH scales. Studies using the NBS scale  
139 were converted to the total scale using the excel macro CO2Sys (Pierrot et al. 2006).  
140 Research examining the effects of ocean acidification were excluded if they did not present  
141 standardized measurements of seawater carbonate chemistry that could allow us to determine  
142 pHT values (Dickson et al. 2007) accurately. That rarely occurred (Smith and Roth 1979;  
143 Borowitzka 1981; Gao et al. 1993). We found a total of 538 calcification rates measured  
144 across different seawater pH treatments that we could include in our analysis.

145 We used the bias-corrected Hedge’s g parameter in our assessment. Hedges g was  
146 calculated as treatment – control divided by the pooled standard deviation. The bias  
147 correction  $\left(1 - \frac{3}{4(df) - 1}\right)$  (Hedges 1981) was used to control for the often-small samples  
148 sizes within studies. There were instances where the total alkalinity anomaly was used to  
149 measure calcification rates. Because some papers only published light calcification rates,  
150 while others also presented dark calcification rates, we excluded all dark calcification  
151 measurements in our analyses.

152 We conducted a meta-analysis in R with the *metafor* package that assessed the effects  
153 of decreasing pH relative to the control values designated by the authors of each study. This  
154 software used a multivariate mixed-effects linear model (function: *rma.mv*), with the unique  
155 observation ID nested within the study ID as random effects. This approach meant that the  
156 impacts of any one study with multiple pH levels were controlled for. For moderators, we  
157 used relative change in  $\text{pH}_T$  between the control and treatment within the experiment ( $\Delta\text{pH}_T$ ;  
158 continuous), the temperature of the experiment ( $^{\circ}\text{C}$ ; continuous), the duration (in days;  
159 continuous), the photosynthetically active radiation level ( $\mu\text{mol}\cdot\text{photons m}^{-2}\text{ s}^{-1}$ ; continuous),  
160 and the taxonomic family of the study organism (categorical: Corallinaceae,  
161 Lithothamniaceae, Lithophyllaceae, Mesophyllumaceae, and Sporolithaceae). In terms of  
162 carbonate chemistry, we choose  $\Delta\text{pH}_T$  because we considered it to be physiologically more  
163 robust than any other seawater carbonate chemistry parameter to test against (Comeau et al.  
164 2018). We additionally tested the model using ‘climate’ (modified from Nybakken 2001  
165 where we also consider the Mediterranean sea as "warm temperate") with tropical, warm-  
166 temperate, cold-temperate and polar, and ‘ocean basin’ (Atlantic, Pacific and Indian Oceans)  
167 as categorical moderators. However, these were subsequently dropped from the model due to  
168 their non-significant effects on the model. Our choice of the taxonomic families for the  
169 moderator followed the current systematics of coralline algae, with reference to Algaebase  
170 (Guiry and Guiry 2021).

171 Cook’s distance (function: *cooks.distance.rma.mv*) was used to test for extreme  
172 outliers, using a conservative cut-off threshold of  $2\sqrt{((k + 1)/(n - k - 1))}$ . This resulted  
173 in a cut-off value of 0.31 for adult responses (removing 12 from 255 observations), and 0.69  
174 for juvenile responses (removing 4 from 57 observations). Studies removed for the adult  
175 responses were from Bergstrom et al. 2020 (6), Fine et al. 2016 (1), Ragazzola et al. 2013 (3)  
176 and Sordo et al. 2018 (2)), and for the juvenile responses were recorded in Page & Diaz-  
177 Pulido 2020 (2), Russell et al. 2011 (1) and Padilla-Gamiño et al. 2016 (1).

178 Meta-analysis results are presented using bubble plots (function: *regplot.rma*). The  
179 results of the multivariate mixed-effects linear model were extrapolated for each of the RCP  
180 2.6, RCP4.5 and RCP8.5 scenarios (function: *predict.rma*) using yearly global mean surface  
181  $\text{pH}_T$  values from the CanESM2 climate model (Canadian Centre for Climate Modelling and  
182 Analysis; CMIP5, ensemble r1i1p1) along with mean values used for the other moderators  
183 (temperature, duration, irradiance and taxonomic family).

184

185 *Proportional responses:*

186 Our initial qualitative review revealed that natural cover of coralline algae along pH/pCO<sub>2</sub>  
187 gradients and the percent cover of juvenile coralline algae on recruitment tiles had been  
188 measured in consistent ways that could easily be compared between studies. We fitted non-  
189 linear models (negative exponential models) for each location individually and pooled the  
190 locations together for an overall model. For both the natural cover and juvenile recruitment,  
191 non-linear models were deemed to provide a better fit compared to linear models based on  
192 lower Akaike information criterion, Bayesian information criterion and estimated standard  
193 error of the residuals. We projected estimates of uncertainty (95 % confidence intervals) in  
194 coverage for the overall models by bootstrapping the model fits for 1,000 runs.

195

196 *Qualitative analysis*

197 Table S1 details the findings of all research we could find on coralline algal responses to OA.  
198 This table summarises what we consider are the main findings. While individual researchers  
199 may argue of the nuances of each study for such a subjective analysis, we used an approach  
200 where four of our author team ultimately reviewed the findings of every study to reduce bias  
201 and misinterpretation. In some instance, we also assessed each study using an alpha = 0.05,  
202 not reduced values used in some research. Arguments of why are beyond this paper, but  
203 corrections for multiple analyses are now considered non-optimal. Especially in this case, i.e.  
204 hundreds of analyses were conducted by the collective research papers, should these be  
205 counted up and corrected for in each study?

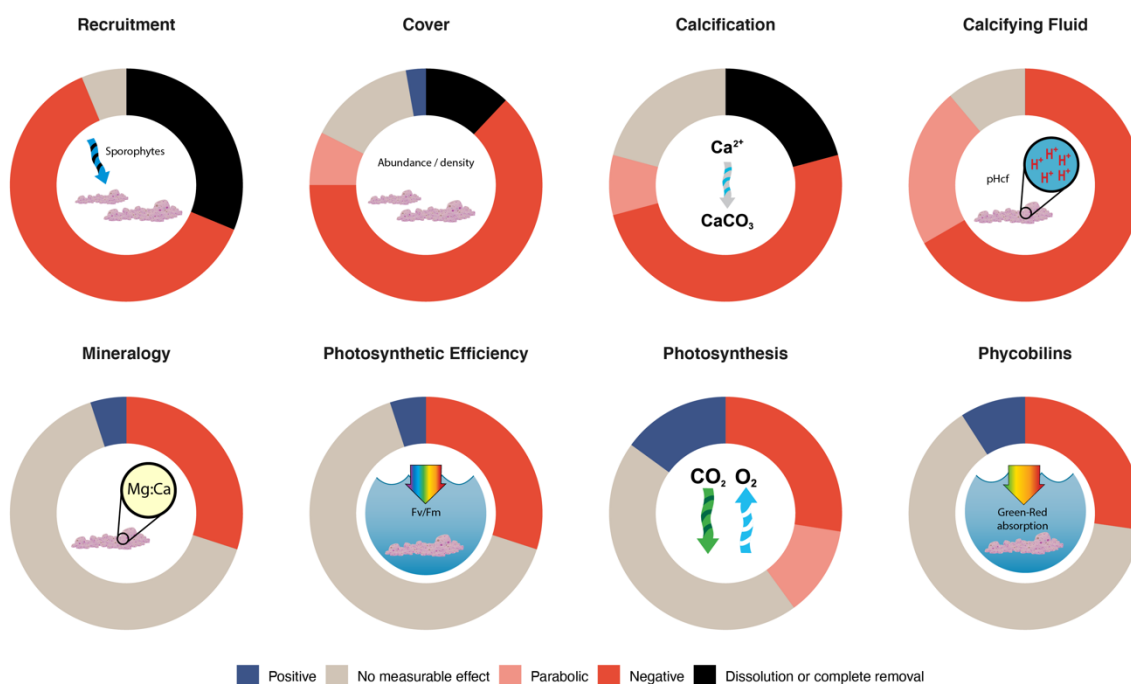
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207 **Results:**208 *Qualitative analysis of main effects*

209 There were 392 recorded responses of coralline algae to the effects of ocean acidification,  
210 within which 108 were responses of calcification/growth. The top 10 most frequently  
211 recorded were as follows with numbers of responses in parentheses: photosynthetic rates (40;  
212 including 26 with respiration rates), cover (24), Fv/Fm (20), mineralogy (20), chlorophyll *a*  
213 content (18), recruitment rates/sizes (16), bleaching/mortality (14), electron transport rates



214 (ETR; 14), phycobilin contents (11) and pH in the calcifying fluid ( $\text{pH}_{\text{cf}}$ ; 9). We focus mostly  
 215 on these parameters for the remainder of this review and in Figure 1. Five of these response  
 216 types were negative more often than not (recruitment 94%, calcification rates 75%, cover  
 217 71% and  $\text{pH}_{\text{cf}}$  67%). Bleaching/mortality was negative 50% of the time; however, this  
 218 response is something visually obvious that was not noted in the majority of laboratory  
 219 studies. Therefore, the proportion of times it truly would have been significantly negative is  
 220 extremely low (< 5%). Negative effects of OA on Fv/Fm, mineralogy (classified as decrease  
 221 in Mg content or increase in aragonite content), photosynthetic rates, chlorophyll *a* content,  
 222 ETR and respiration occurred 30% or less in occasions where they were recorded. Positive  
 223 responses of photosynthetic rates (15%) and ETR (14%) occurred more frequently for these  
 224 top ten measured metrics than any others. Parabolic responses were noted in 21 responses  
 225 from all measured 392 responses (i.e. 5% of times).

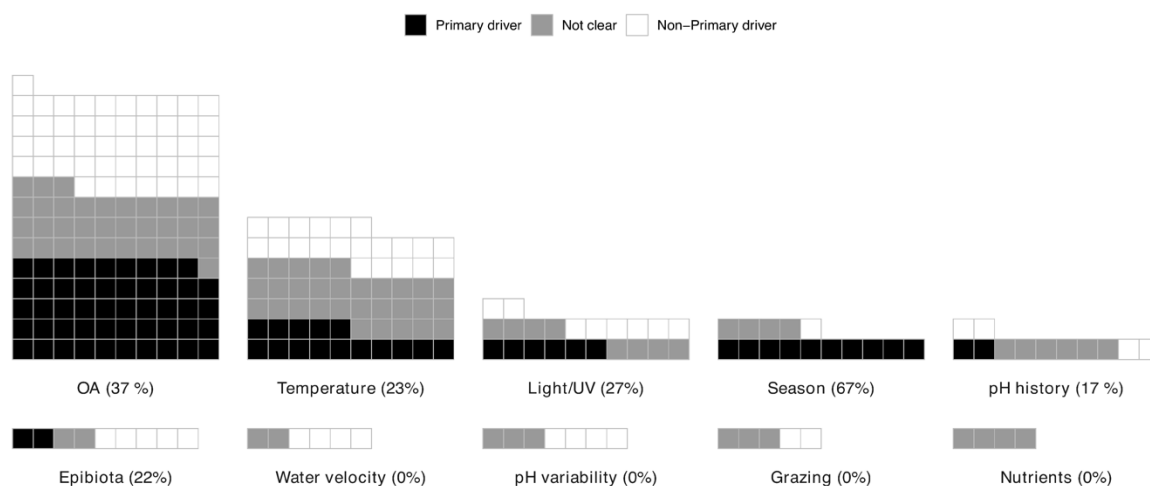


227 Figure 1: Effects of OA on coralline algae recorded in the literature, with the eight most  
 228 negatively affected response parameters shown here for those parameters that were measured  
 229 at least nine times. Responses are described as either 'positive', 'no measurable effect',  
 230 'parabolic', 'negative' or 'dissolution or complete removal', see methods for more details.

231

232 *Dominant driver*

233 Seasonality was the dominant driver more times than any other (67%: Figure 2). OA was the  
 234 next, being the putative dominant driver 37% of occasions, then light quality/quantity (27%),  
 235 temperature (23%), epiphyte/turf presence (22%), and past pH history of the coralline algae  
 236 (17%). Water velocity, pH variability, herbivory/grazing and nutrients were never dominant.



238 Figure 2: Number and proportion of times different drivers were assessed in combination  
 239 with OA or another driver in 3+ factorial approaches. In parentheses are the percentage of  
 240 times the driver was dominant. See methods for the definition of dominant.

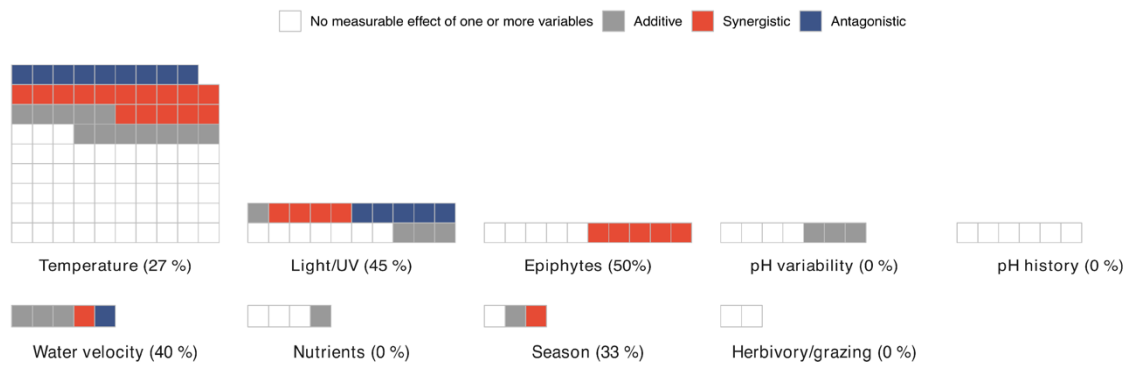
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## 242 *Interaction types*

243 Here we split the interactions by the measured response type (e.g. calcification rates, etc.) and  
 244 the driver that was tested. We report response types that were measured in multi-driver  
 245 experiments 10 or more times. Photosynthetic rates had the highest number of “non-  
 246 surprising” interactions (additive or unmeasurable main effects and interactions; 87%),  
 247 followed by Fv/Fm (83%), calcification rates (72%) and respiration rates (60%). There was a  
 248 total of 21 antagonistic and 18 synergistic response from 153 recorded, 14% and 12%  
 249 respectively.

250 There were clear patterns in the driver types that more often resulted in these  
 251 antagonist and synergistic interactions. These were mostly drivers that can alter the pH at the  
 252 surface of coralline algae or their photo-physiology: epiphyte/turf presence (50% of  
 253 antagonistic + synergist interactions), light (45%) and water velocity (40%). Season (33%)  
 254 and temperature (27%) also had some antagonistic or synergistic interactions. Interactions

255 with pH variability, nutrients, herbivory and past pH history did not result in synergistic or  
 256 antagonistic effects.



258 Figure 3: Number of different interaction types between OA and other drivers within  
 259 coralline algal OA research. The proportion of synergistic and antagonistic interactions is  
 260 indicated within parentheses. Note here we define “no measurable effect of one or more  
 261 variables” as being separate from “additive”, with the former being occasions where one or  
 262 both drivers have no measurable effect in either direction, and the addition of the other driver  
 263 does not change the effect of OA.

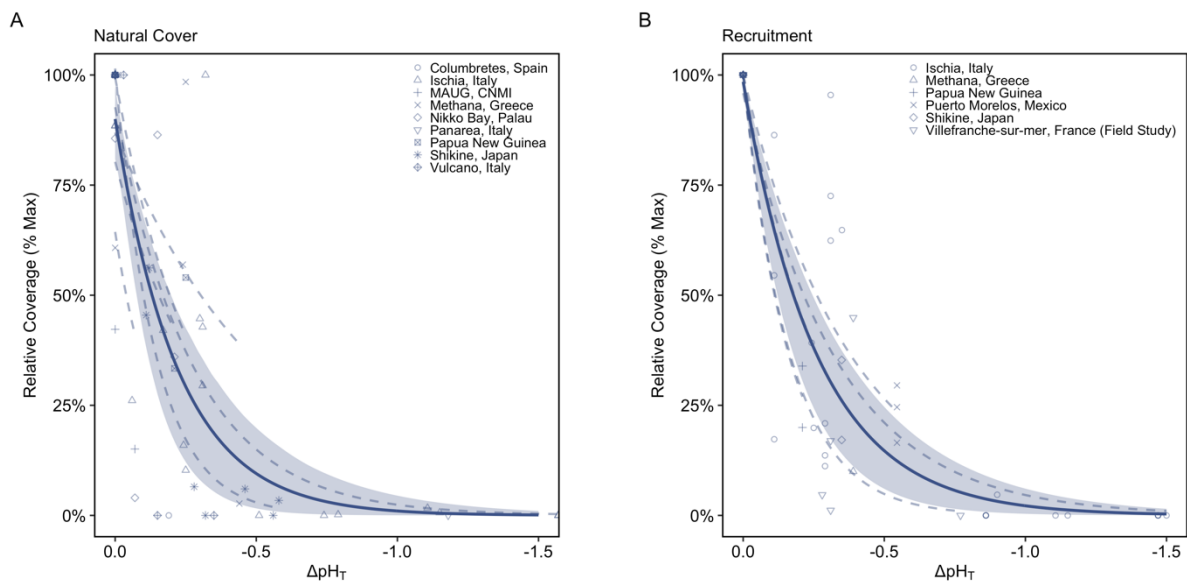
264

265 *Proportional effects*

266 *Field cover and recruitment rates*

267 There were large reductions in coralline algae cover as pH declined, with the non-linear  
 268 model predicting that relative coverage may drop to ~ 15 % (2 – 27, 95 % CI) of their  
 269 associated reference/control site(s) once  $pH_T$  declines below -0.4, and continue to approach  
 270 towards zero covers (Figure 4a). Recruitment cover demonstrated a similar decline as pH is  
 271 reduced, albeit less steeply, with ~ 21 % (11 – 32, 95 % CI) compared to the

272 reference/control site(s) at a  $\Delta\text{pH}_T$  decline of -0.4 (Figure 4b).



273

274 Figure 4: A) Relative cover change (%) of coralline algae at reduced pH sites in the field  
 275 compared to nearby control sites. B) Relative coverage (%) of coralline algal recruits on  
 276 settlement tiles at reduced pH locations compared to nearby controls.

277

## 278 *Meta-analysis*

279  $\Delta\text{pH}_T$  had a significant relationship with calcification in the multivariate meta-analysis model  
 280 for adult coralline algae (Fig. 5), indicating that as  $\text{pH}_T$  is reduced, the negative impact on  
 281 calcification may become greater. Based on the model, after a threshold of  $-0.062 \Delta\text{pH}_T$  (at  
 282 Hedge's  $g$ :  $-0.53 \pm 0.53$  95% CI), adult coralline algae had their calcification significantly  
 283 reduced, reaching (for example) effect size of  $-1.68 \pm 0.36$  (95 % CI) at  $-0.4 \Delta\text{pH}_T$  (where  
 284 the effect size of Hedge's is deemed to be a 'large effect' at 0.8). Some phylogenetic  
 285 differences were found within the adult coralline algae responses, where those coralline algae  
 286 from the family Mesophyllumaceae had a significantly different slope than the other families  
 287 (Table S2, S1).

288 There were fewer responses available in the literature for juvenile coralline algae, and  
 289 those that were available assessed growth/calcification over a smaller range of  $\Delta\text{pH}_T$ , which  
 290 likely led to no significant effect of  $\Delta\text{pH}_T$  on the model outcome (Table S3, Fig. S2).

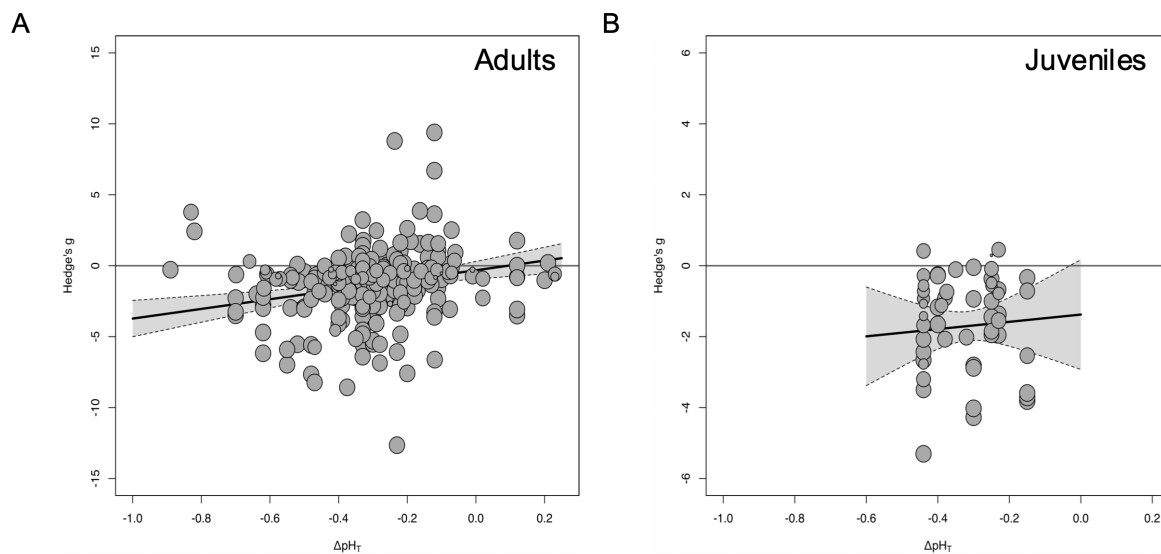
291 Although the calcification response did not significantly change with  $\Delta\text{pH}_T$ , the observations  
 292 themselves were still almost entirely negative; for example, at  $-0.4 \Delta\text{pH}_T$  an effect size of

293  $-1.79 \pm 0.57$  (95 % CI) was observed, indicating a significant and large negative effect on the  
 294 calcification of juvenile coralline algae. In addition, there was a significant difference  
 295 between the responses of juveniles from both families Lithophyllaceae and Corallinaceae  
 296 compared to other families. This is likely simply because there is sufficient statistical power  
 297 to assess differences between these two families but not enough data for the others.

298 When considering our analysis in the context of the RCP scenarios, we found the  
 299 effects of altered seawater pH were highly dependent on the RCP scenario being considered.  
 300 For the subtle effects of seawater pH under conditions representing RCP2.6, we predict non-  
 301 significant declines in adult calcification by 2100 compared to the present-day (Figure 6). We  
 302 also predict significant declines in calcification under RCP4.5 and 8.5 by 2100 compared to  
 303 present-day, with calcification rates declining drastically in the RCP8.5 scenario as this  
 304 scenario deviates from RCP4.5 around the year 2040 (Figure 6).

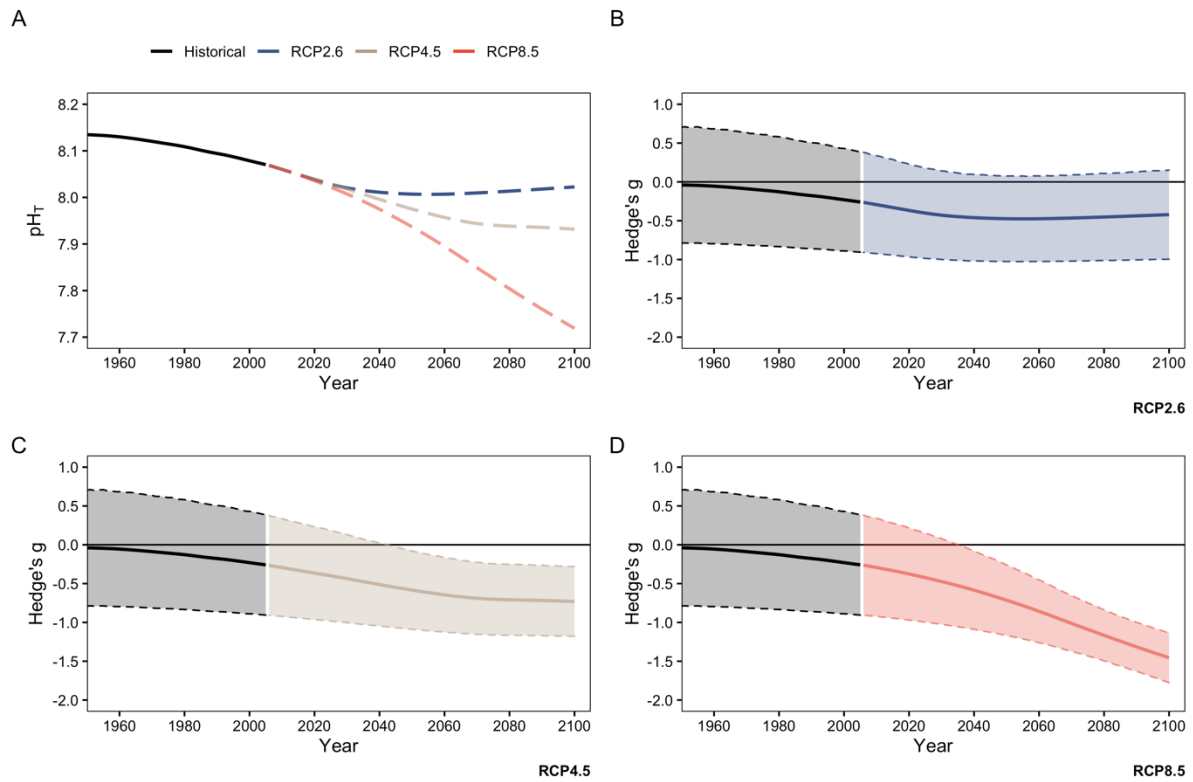
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306



308 Figure 5: (A and B) – Bubble plots for adults and juveniles, respectively, showing how the  
 309 effect size on calcification changes with  $\Delta pHT$  (while still considering the other predictors).  
 310 The size of the points indicates the weighting, and the dashed line indicates 95% CI.

311



313 Figure 6: (A) Predicted global mean  $pH_T$  level historically, and projected for RCP2.6, 4.5 and  
 314 8.5 scenarios up to the year 2100 using the CanESM2 climate model. (B – D) The mean  
 315 effect size  $\pm$  95% CI for adults at the predicted global mean  $pH_T$  level historically, and for  
 316 RCP2.6 (B), 4.5 (C) and 8.5 (D) scenarios up to the year 2100.

317 **Discussion:**

318 We demonstrate here that OA is a major threat to the persistence of coralline algae in the  
319 future using multiple lines of evidence. Early studies that reported severe declines in coralline  
320 algal cover and recruitment at high CO<sub>2</sub> sites (Hall-Spencer et al. 2008; Martin et al. 2008;  
321 Fabricius et al. 2011), and declines in calcification and recruitment (Anthony et al. 2008;  
322 Jokiel et al. 2008; Kuffner et al. 2008; Martin and Gattuso 2009; Russell et al. 2009) have  
323 been fully corroborated by a host of research demonstrating that such effects are the norm,  
324 rather than the exception. Though not explicitly tested, no such “decline effect” (Clements et  
325 al. 2020), as has been purported in other taxa, was found here. Indeed, the studies that  
326 showed the larger effect sizes, were some of the most recent (Barner et al. 2018; Sordo et al.  
327 2018; Qui-Minet et al. 2019; Bergstrom et al. 2020). Declines in calcification, recruitment,  
328 coralline cover at naturally high CO<sub>2</sub> analogues (and less commonly, pH within the  
329 calcification fluid (pH<sub>cf</sub>)) were the norm, rather than the exception across all published  
330 literature.

331 Despite consistent responses of many of these vital ecological and physiological  
332 processes, some other physiological processes had inconsistent responses to OA. By  
333 understanding such distinctions, we can better assess how coralline algal communities are  
334 likely to be altered by future climate change. For example, some Authors showed that OA  
335 would cause widespread coralline algal mortality due to bleaching and declines in pigment  
336 concentrations and decreases in photosynthetic rates or DIC uptake (Anthony et al. 2008;  
337 Martin and Gattuso 2009). These results led to speculation that OA could cause similarly  
338 catastrophic impacts for coralline algal dominated reefs as the increased temperature does for  
339 corals. Fortunately, we find little evidence for this. However, the declines in calcification  
340 observed here still threaten the ability of coralline algae to perform vital ecological services,  
341 such as providing settlement substrata or building and cementing biogenic reefs (Cornwall et  
342 al. 2021). Additionally, we find inconsistent evidence for other benefits to photosynthesis,  
343 photo-physiology or DIC uptake that earlier studies proposed (Cornwall et al. 2012), or  
344 changes in mineralogy that could either benefit or impair their physiology (Ries 2011; Diaz-  
345 Pulido et al. 2014). That highlights the importance of pooling studies for reanalysis to refine  
346 hypotheses. Conversely, it was clear that OA usually impacts pH<sub>cf</sub> of the coralline algae (e.g.  
347 Cornwall et al. 2017; Donald et al. 2017), though other elemental or geochemical differences  
348 were unclear. Irrespective of their findings, these earlier studies collectively led and

349 developed the field of coralline algal responses to OA, allowing the current synthesis of the  
350 results.

351

352         The relative importance of global and local drivers in determining ecological change  
353 is highly dependent on the taxa being assessed. For corallines, OA was the dominant driver  
354 more often than any other local or global stressor it was assessed against, with seasonality  
355 being the only driver that was dominant more often. Despite the sensitivity of many taxa to  
356 changes in seawater temperature, it was rarely the clear dominant driver of coralline algal  
357 responses. This observation supports recent work questioning the relative importance of  
358 temperature on coralline algal physiology, compared to more temperature-sensitive  
359 organisms such as corals or kelps (Cornwall et al. 2019; Anton et al. 2020). The combined  
360 effects of ocean warming and OA are better constrained in the literature than the role of other  
361 drivers. However, there is still very little known regarding its impacts (Cornwall et al. 2019),  
362 and it should be noted that multiple responses recorded here were from the same study (e.g.  
363 multiple species or multiple response metrics). While 15 “responses” were recorded where  
364 temperature played a dominant role, 6 of these were recorded from the same study (Vásquez-  
365 Elizondo and Enríquez 2016). Conversely, there is little understanding of how marine  
366 heatwaves will impact coralline algae, and these would be expected to have much greater  
367 impacts than slow ongoing ocean warming that was often simulated in most research we  
368 assessed here, for example that possessed treatments that were simply 2 or 3 degrees above a  
369 control seawater. Likewise, the role of light was often complex and only sometimes (6/22)  
370 played a dominant role in determining the physiological state of the coralline algae. It is more  
371 difficult to synthesise the effects of light quality and quantity into a single category to assess  
372 their role as dominant or not. Exposure to differing spectra or total PAR daily doses (either  
373 too low or too high) within the realms of that experienced at any one collection site could  
374 elicit stronger responses than changes in light. Multi-stressor experiments need to carefully  
375 choose stress levels and combinations to obtain the most informative results (Boyd et al.  
376 2018).

377         Here we found that environmental/ecological factors that were most likely to elicit  
378 antagonistic or synergistic effects were those that impacted either photo-physiology or pH at  
379 the surface of coralline algae. A possible reason for that may be that the pH experienced by  
380 the organism is more important than that of the bulk seawater. At the surface of most

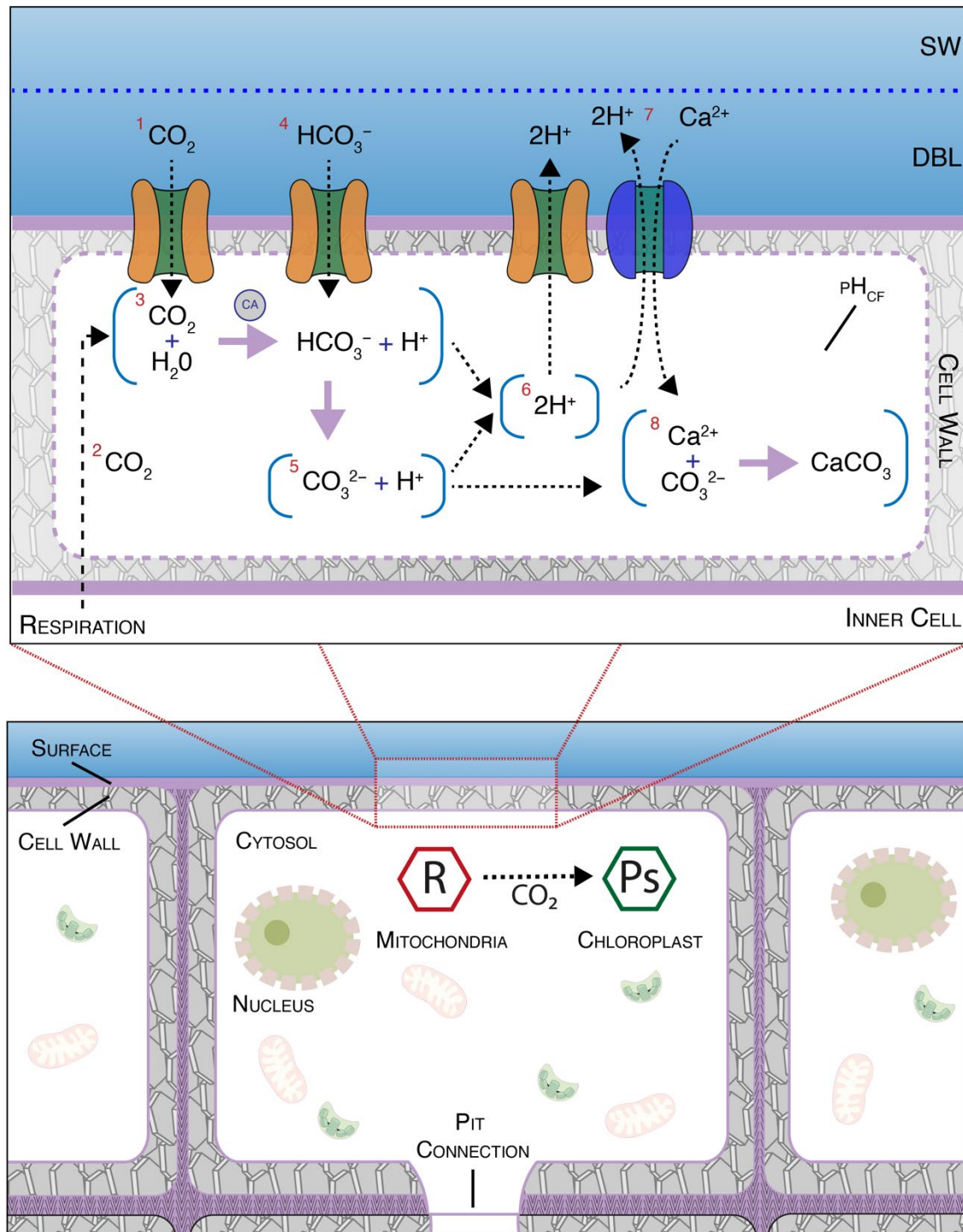


381 photoautotrophs, pH is set by the thickness of the diffusion boundary layer (DBL: a region of  
382 seawater where the movement of dissolved substances is primarily by molecular diffusion).  
383 The theoretical maximum of the DBL is set by water velocity near the organism, and this  
384 near-surface velocity would be reduced by the presence of any epiphytes or canopy-forming  
385 species above the coralline algae, hence increasing pH gradients (Cornwall et al. 2015). The  
386 pH achieved within the DBL will be dependent on the rate of metabolic activity within it (e.g.  
387 photosynthesis or respiration) and the DBL thickness. Therefore, it is not surprising that  
388 drivers that modify pH within the DBL (light, epiphytes, and water velocity) often nullify the  
389 impacts of OA (Cornwall et al. 2014; Guy-Haim et al. 2020), although not always when  
390 nutrient concentrations are low (Comeau et al. 2014; Comeau et al. 2019b). However, the  
391 mechanisms responsible for the antagonistic roles of season and temperature are harder to  
392 disentangle. Indeed, interpreting many of the seasonal impacts on coralline algae were  
393 difficult during the process of constructing Table S1, especially when additional factors such  
394 as temperature and light were either modified at each season or tested interactively (e.g. Qui-  
395 Minet et al. 2018; Legrand et al. 2019; Qui-Minet et al. 2019).

396 Changes to the ecological role of coralline algae in a future low pH/high CO<sub>2</sub> ocean is  
397 uncertain; we demonstrate their ability to calcify, recruit and persist in abundances similar to  
398 those today are impacted by OA. We also demonstrate that their calcification rates will depart  
399 significantly from that of today's as OA intensifies. The effects noted in our qualitative  
400 analysis are mostly linear, with few parabolic or rapid "tipping point" type responses. Our  
401 meta-analyses indicate that under RCP8.5 they might be particularly severely impacted,  
402 highlighting the urgent need to reduce our CO<sub>2</sub> emissions to allow these critical species  
403 continuing their role as reef accrete and preferred settlement substrate for many invertebrate  
404 species (Fabricius et al. 2017). Surface seawater pH has been declining since the industrial  
405 revolution, meaning that the impacts of OA have likely already manifested in some way  
406 across global oceans (McCoy and Pfister 2014), and this is supported by comparing sites with  
407 naturally low, pre-industrial levels of seawater CO<sub>2</sub> (Agostini et al. 2018). Select coralline  
408 algae can acclimatise over multiple generations to completely resist the effects of OA  
409 (Cornwall et al. 2020; Moore et al. 2021). However, whether all species can respond  
410 similarly, if the presence of rapidly growing competitors reduces this ability, or whether there  
411 are trade-offs in genetic diversity that occur in these populations remains unknown.

412 The negative effects on calcification manifest across all major oceans and across the  
413 range from tropical to polar climates. Specific phylogenies appeared more robust than others,

414 particularly the adult Mesophyllumaceae. A scientific effort in monitoring the abundance or  
415 ecology of coralline algae beyond simply acknowledging their collective presence or absence  
416 is lacking. That is hindered by the only recent understanding that coralline algae diversity is  
417 greater than previously believed (Gabrielson et al. 2018; Twist et al. 2019; Peña et al. 2021b),  
418 and extends far beyond the previous mostly morphological groupings and has important  
419 implications for much of the previous work in this field (Twist et al. 2020). Recent molecular  
420 work shows that some groups of coralline algae are more vulnerable to OA than others and  
421 that the projected adverse effects on coralline algal biodiversity have been greatly  
422 underestimated (Peña et al. 2021a). Likely changes in species abundances, competitive  
423 hierarchies, and physiologies (McCoy and Pfister 2014; McCoy and Ragazzola 2014) before  
424 the advent of molecular technology make it challenging to establish a baseline abundance for  
425 most coralline algae. Future research should focus on pairing physiological and molecular  
426 identifications wherever possible.



428 Figure 7: Schematic showing movement of dissolved substances in and out of coralline algal  
 429 cells. This gives an overview of the processes that could be affected by ocean acidification, it  
 430 is not intended to describe the full calcification process. In some taxa there can be multiple  
 431 cells between seawater and the site of calcification hereafter “Calcifying fluid = CF”.  
 432 Provision of photosynthesis by 1) movement of  $\text{CO}_2$  (or external carbonic anhydrase  
 433 mediated conversion of  $\text{HCO}_3^-$  to  $\text{CO}_2$ ) across aquaporins or through seawater entering the

434 cell wall and 2) produced during respiration. 3) The reaction from  $\text{CO}_2$  to  $\text{HCO}_3^-$  will be  
435 catalysed by internal carbonic anhydrase to some extent within the cell, while 4) active  
436 transport of incoming  $\text{HCO}_3^-$  will occur via  $\text{CO}_2$  concentrating mechanisms involving influx  
437 or efflux to balance the charge (not shown). 6) Increasing  $\text{pH}_{\text{cf}}$  will be due to proton pumps or  
438 7)  $\text{Ca}^{2+}$  ATPases or other mechanisms of  $\text{H}^+$  efflux. 8) Elevated saturation state of  $\text{CaCO}_3$   
439 will increase precipitation rates within the calcifying fluid pockets.

440

441 The physiology of coralline algae is poorly understood, hindering our ability to  
442 interpret many of the observed responses to seawater carbonate chemistry. Their calcification  
443 is impacted by declining seawater pH via both increased dissolution rates and an inhibition in  
444 the precipitation of calcium carbonate. The amount of dead exposed skeletal material within  
445 any experiment would therefore strongly influence the rate of dissolution, further impairing  
446 our ability to compare experimental results. After decades of models based on eco-  
447 physiology, scanning electron microscope images, and some carbon isotope work, new  
448 models and techniques have emerged. These have allowed us to determine that calcification  
449 in coralline algae is highly controlled by the organism, while at the same time is still dictated  
450 by environmental conditions. Recent models have demonstrated that control within the site of  
451 calcification, likely  $\mu\text{m}$  or smaller sized pockets of instantaneous fluid that we term the  
452 “calcifying fluid” (CF), is strongly influenced by ocean acidification. The  $\delta^{11}\text{B}$  of calcium  
453 carbonate tracks the pH of the solution in which it is formed (Klochko et al. 2006). For  
454 coralline algal calcite, experimental work has demonstrated that this proxy tracks pH within  
455 the CF, finding that it is highly upregulated, beyond what is possible simply due to  
456 photosynthetic activity, in the majority (but not all) of the species measured to date (Cornwall  
457 et al. 2017; Donald et al. 2017; Cornwall et al. 2018; Anagnostou et al. 2019; Comeau et al.  
458 2019a; Comeau et al. 2019b). Importantly, the species-specific magnitude of its decline under  
459 simulated ocean acidification is linked with a greater capacity for calcification to resist ocean  
460 acidification for some species (Cornwall et al. 2017). However, it is likely to be more  
461 complex than suggested by those initial observations

462 Both DIC and the saturation state of calcite within CF are likely important factors in  
463 the regulation of calcification in coralline algae, even if these have not been measured  
464 directly with existing technology. Raman spectroscopy (FWHM) and B/Ca ratios track  $\Omega_{\text{cf}}$   
465 and  $\text{DIC}_{\text{cf}}$  in coral aragonite (Holcomb et al. 2016; DeCarlo et al. 2017; McCulloch et al.

2017). However, similar calibrations do not exist for coralline algal high Mg calcite. In both corals and coralline algae these parameters are altered by the same environmental factors (pH, DIC, light and water flow) in the same direction (with some species-specific offsets), leading to the more likely scenario that they also represent  $\Omega_{cf}$  and  $DIC_{cf}$  in coralline algal calcite. There have been some suggestions for corals that B/Ca may even represent a differential uptake of boron under changing DIC (Gagnon et al. 2021). However, this would not account for the remarkably similar  $DIC_{cf}$  estimated with electrodes (Cai et al. 2016). Overall, these proxies would indicate that coralline algal calcite saturation state and DIC is likely much higher than that of seawater. This would rule out photosynthetic control over carbonate chemistry within the CF, even if DIC was slightly higher than that of seawater, as DIC would be expected to decrease under higher photosynthetic rates. However, that is not to say that higher photosynthetic rates would not further elevate  $pH_{cf}$  or  $\Omega_{cf}$  if this was in concert with increased pumping of  $H^+/Ca^{2+}$  due to increasing energy. This potential  $\Omega_{cf}$  remains invariant under many conditions, though it is influenced by Mg content and should be adjusted accordingly (Perrin et al. 2016). There are, however, two conditions in which it is altered. Like in corals, it appears that very rapid calcification is associated with lower  $\Omega_{cf}$  values for some fast growing juveniles (Cornwall et al. 2020); which could be associated with the drawdown of  $Ca^{2+}$  and would slow the precipitation of calcium carbonate when combined with lower  $pH_{cf}$  under OA. Conversely, this could also represent changes in  $\Omega_{cf}$  driven by an inability to adequately regulate a combination of  $DIC_{cf}$  and  $pH_{cf}$ .  $\Omega_{cf}$  declined in *Neogoniolithon* sp. (but not *Sporolithon durum*) when seawater DIC was reduced experimentally (Comeau et al. 2018). Interestingly, *S. durum*  $pH_{cf}$  increased and  $DIC_{cf}$  declined in response to the same conditions, with  $pH_{cf}$  declining and  $DIC_{cf}$  increasing under lower seawater pH, but with no effect on *Neogoniolithon* sp. Low light can reduce  $\Omega_{cf}$  (Comeau et al. 2019b), likely due to a reduced ability to pump  $Ca^{2+}$ , as  $pH_{cf}$  and  $DIC_{cf}$  remain constant under low light. Further species-specific investigations are required, especially those that assess apparent phylogenetic differences in calcification modes (Auer and Piller 2020) with differential responses of CF chemistry. Though previously identified morphological differences in calcification modes (Auer and Piller 2020) did not match more resistant or susceptible taxa here, it could reveal important information if indeed crystal orientation does impact the CF chemistry or responses to OA and other abiotic factors.

There are still many fundamental questions left unanswered in understanding the responses of coralline algae to OA. Coralline algal CF chemistry is species-specific and has

499 large uncertainties compared to that of aragonitic corals. While small windows of knowledge  
500 have been opened by using these proxies, there is still much work to be done in calibrating  
501 B/Ca and FWHM to understand the actual values of  $\text{DIC}_{\text{cf}}$  and  $\Omega_{\text{cf}}$  in coralline algae and how  
502 they respond to different environmental drivers, particularly OA. Taxonomic variation in  
503 responses, and its influence on other factors that we explored here can only be guessed at  
504 currently with the existing issues of non-molecular identification in the majority of studies  
505 included here (Twist et al. 2020). Better integration of molecular work while assessing their  
506 response to OA is needed in the laboratory. How this can be implemented in field settings  
507 such as natural  $\text{CO}_2$  vents is a more difficult consideration. If plans can be implemented to  
508 include molecular identification in  $\text{CO}_2$  vent work, this will go a large way in disentangling  
509 which species are tolerant there, as has been recently completed at two  $\text{CO}_2$  seep sites (Peña  
510 et al. 2021a). Additionally, the next steps are to combine “omics” tools with physiological  
511 and geochemical tools to better determine the mechanisms responsible for individual species  
512 tolerance to OA, and how and why these tolerances can change over multiple lifetimes in  
513 some coralline algae (Cornwall et al. 2020). Here, we do not attempt to fully discuss  
514 calcification mechanisms, but future inclusions of emerging techniques with morphological  
515 knowledge (e.g. Auer and Piller 2020) will also further our ability to understand taxa-specific  
516 responses. Only by not ignoring all available information will progress be made in better  
517 predicting the future of ecosystems in which coralline algae act as important foundation  
518 species.

519

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